

Upper Llandovery to middle Wenlock (Silurian) lithostratigraphy and chitinozoan biostratigraphy in southwestern Estonia and northernmost Latvia

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Abstract. The lithostratigraphy and chitinozoan biostratigraphy of the Adavere, Jaani, and Jaagarahu regional stages (upper Llandovery to middle Wenlock) in the Ikla (Estonia) and Staicele (Latvia) drill cores were studied. Most of the upper Aeronian to lower Homerian local chitinozoan biozones are represented in these cores, except Interzone III, *Angochitina longicollis*, *Conochitina proboscifera*, and *Margachitina margaritana* biozones, partly corresponding to barren red beds of the Velise Formation, partly to a stratigraphical gap at the Llandovery–Wenlock transition. The presence of *Stricklandia laevis* shows that the Rumba Formation of the Adavere Stage belongs to the Telychian. The succeeding, Velise Formation has a reduced thickness and is represented only by its lower portion. Three stratigraphical gaps were established at the boundaries of the regional stages. The hiatuses at the junctions of the Raikküla–Adavere and Adavere–Jaani stages have a considerable areal extent and stratigraphical range. Probably, they reflect global glacio-eustatic falls of sea level at the end of the Aeronian and Telychian corresponding to the glaciations in South America. A gap at the base of the Jaagarahu Stage was recognized only in the Ikla core and has local significance.

Key words: lithostratigraphy, chitinozoan biostratigraphy, stratigraphical gaps, Llandovery, Wenlock, Estonia.

INTRODUCTION

The subsurface sequence of Llandovery carbonate and calcareous-argillaceous rocks is thickest and most complete in southwestern mainland Estonia and northernmost Latvia. The Llandovery sequence reaches 231 m thick in the Ikla drill core (Estonia) and 222 m in the Staicele drill core (Latvia) (Fig. 1), more than twice the thickness of the series in outcrops in central and western Estonia. The Ikla and Staicele sections exhibit the transition from shelly to graptolite biofacies and have received attention as the type sections of several lithostrati-



Fig. 1. Location of drill holes in southwestern mainland Estonia and northernmost Latvia figured in Figs. 2–4.

graphical units (Saarde Formation, Ikla, Lemme, and Staicele members). However, up to now only the more calcareous rocks of the Juuru, Raikküla, and the lower part of the Adavere regional stages, corresponding to the lower and middle Llandovery, have been sufficiently well studied. The litho- and biostratigraphy of the Ikla and Staicele cores have been treated by Kaljo & Vingisaar (1969), Ulst (1973), and Gailite et al. (1987). The data from the Ikla core have been used for elaboration of the regional chitinozoan biozonation (V. Nestor 1976, 1994) and as a constituent part of several geological transects (e.g. H. Nestor & Einasto 1982; Einasto 1984; H. Nestor 1997).

Only scanty published data are available on the upper part of the Ikla and Staicele cores which contain argillaceous rocks of the upper Llandovery (Velise Formation) and the lower–middle Wenlock (Jaani and Jaagarahu regional stages). Ulst (1973) considered that the whole complex of the uppermost marlstones in the Staicele core and a great part of the Ikla core belong to the Velise Formation of the Adavere Stage. Aaloe (1970) supposed that the Ikla core ends at the top of the Jaani Stage. Recent study of chitinozoans from the upper portions of the Ikla and Staicele cores proved that both reach much higher in the stratigraphical succession than previously supposed.

The present paper is based on the macrolithological investigation and sampling of the Staicele core at the Katlakalns depository of the Latvian Geological Survey near Riga and the Ikla core at the Keila depository of the Estonian Geological Survey near Tallinn. Also, unpublished descriptions of the Ikla core by Peeter Vingisaar and Rein Einasto were used.

The lithostratigraphy and chitinozoan biostratigraphy of the Adavere, Jaani, and Jaagarahu regional stages are considered. The Adavere Stage corresponds

roughly to the upper Llandovery (Telychian), the others to the lower and middle Wenlock (Sheinwoodian to lower Homerian). The lithostratigraphical units used are the same as those in the recent Silurian stratigraphical scheme of Estonia (H. Nestor 1997).

LITHOSTRATIGRAPHY

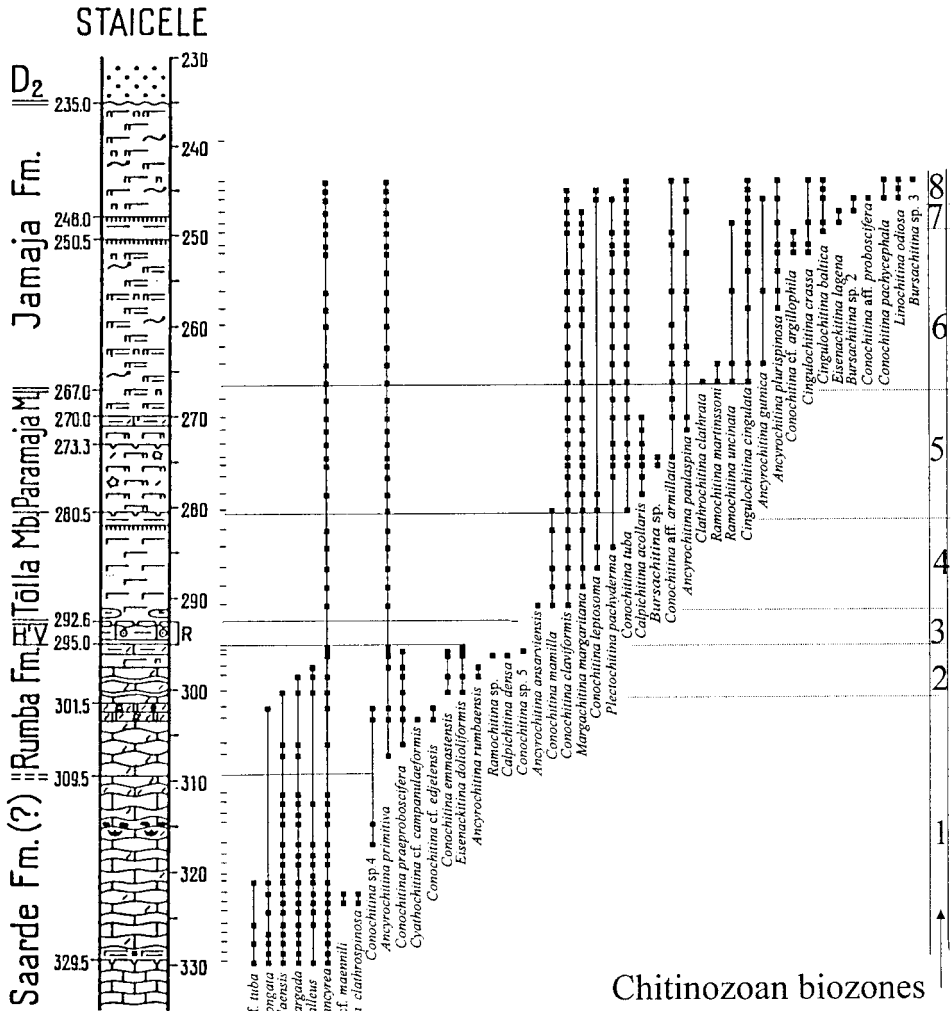
Adavere Stage

The Adavere Stage is represented in the Staicele and Ikla cores by the Rumba (below) and Velise (above) formations (Figs. 2 and 3). In western Estonia the **Rumba Formation** typically consists of variably argillaceous, nodular to wavy-bedded biomicritic limestones (packstone and wackestone) with accumulations of the brachiopod *Pentamerus oblongus* at certain levels. In general, the formation is characterized by distinct low-rank cyclicity (Einasto et al. 1972). In the Ikla and Staicele cores the cyclicity is less conspicuous and pentamerids are rare, representing a mixed *Pentamerus*–*Stricklandia* community. From the Rumba Formation of the Ikla core, Rubel (1977) has identified *Stricklandia lens ultima* (= *Stricklandia laevis* (Sowerby)), which is a guide fossil of the Telychian.

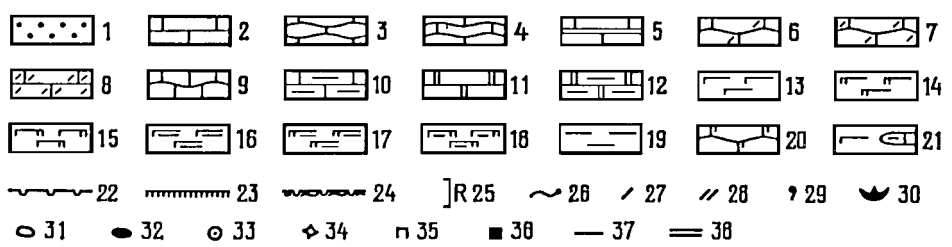
In the interval of 295.0–329.5 m of the Staicele core the Rumba-type nodular wackestones are almost twice as thick as the Rumba Formation in the Ikla core (304.8–322.3 m). In previous publications (Ust 1973; Gailite et al. 1987) the whole or a great part of these strata is included in the Staicele Member of the Saarde Formation. However, from the lithostratigraphical point of view it seems plausible to include the entire complex of nodular wackestones in the Rumba Formation, assuming that the formation of the Rumba-type rock might have commenced somewhat earlier in the vicinity of Staicele.

The **Velise Formation** has a reduced thickness in both cores and bears some specific lithological features. In the Staicele core the interval of 292.5–295.0 m is represented by interbedded reddish-brown to varicoloured, bioturbated marlstones and argillaceous limestones containing an interlayer with ferruginous (iron) oolites at a depth of 293.6–293.7 m. The topmost 15 cm of the interval contains stromatolites and several limonitic encrustations, probably representing discontinuity surfaces.

In the Ikla core the Velise Formation ranges through the interval of 296.0–304.8 m. It is represented by reddish-brown to grey, argillaceous marlstone with limestone nodules. Some bentonite interlayers, discontinuity surfaces, and scattered oolites are observed in the middle of the interval. At the base of the formation there occurs a series of at least eight limonitic discontinuity surfaces. Several discontinuity surfaces are recorded also higher up in the section, one lying at the top of the formation. Rounded pebbles of pyritized argillaceous limestone occur above the discontinuity surfaces at depths of 297.1 and 298.2 m. The Velise Formation is almost barren of micro- and macrofossils. The reduced thickness of rocks, restricted number of ash layers, presence of oolites, discontinuity surfaces



- Conochitina* aff. *tuba*
- Conochitina* *elongata*
- Conochitina* *iklaensis*
- Conochitina* *alargada*
- Conochitina* *mallevs*
- Ancyrochitina* *ancysrea*
- Spraachitina* cf. *maupinili*
- Ancyrochitina* *clathrata*
- Conochitina* sp. 4
- Ancyrochitina* *primitiva*
- Conochitina* *praeroboscifera*
- Cyathochitina* cf. *campanulaeformis*
- Conochitina* cf. *edtelensis*
- Conochitina* *emmastensis*
- Eisenackitina* *dololiformis*
- Ancyrochitina* *rumbaensis*
- Ramochitina* sp.
- Calpiclitina* *densa*
- Conochitina* sp. 5
- Ancyrochitina* *ansarvensis*
- Conochitina* *mamilla*
- Conochitina* *claviformis*
- Margachitina* *margariana*
- Conochitina* *lepuscula*
- Plectochitina* *pachyverna*
- Conochitina* *tuba*
- Calpiclitina* *acollaris* sp.
- Bursachitina* sp.
- Conochitina* aff. *armillata*
- Ancyrochitina* *naulaspina*
- Clathrochitina* *clathrata*
- Ramochitina* *marinssoni*
- Ramochitina* *unicata*
- Cingulochitina* *cingulata*
- Ancyrochitina* *gunica*
- Ancyrochitina* *plurispinosa*
- Conochitina* cf. *argillophila*
- Cingulochitina* *crassa*
- Cingulochitina* *ballica*
- Eisenackitina* *lagena*
- Bursachitina* sp. 2
- Conochitina* aff. *proboscifera*
- Conochitina* *pachycephala*
- Linochitina* *odiosa*
- Bursachitina* sp. 3



with lithoclasts, and layered stromatolites indicate that the sequence is probably incomplete. The reddish colour of the rocks suggests that only the lower part of the Velise Formation is preserved in the Ikla and Staicele cores, because in several adjacent, stratigraphically more complete cores (Kihnu, Varbla, Seliste, etc.) only the lowermost part of the Velise Formation is red.

Jaani Stage

The Jaani Stage is represented in the considered sections by the stratigraphical units characteristic of the South Estonian–North Latvian Confacies Belt. The lower part of the stage contains calcareous (in Staicele) to argillaceous (in Ikla) dolomitic marlstones conventionally attributed to the **Tõlla Member** of the **Riga Formation** (Aaloe 1970). However, intercalations of dark grey calcareous mudstone with graptolites are lacking in these marlstones and therefore the sections have an intermediate position between the Riga and Jaani formations.

The upper part of the Jaani Stage is represented by marlstones of the **Paramaja Member** of the **Jaani Formation**. In the Staicele core the rock is more calcareous and contains much bioclastic material, but the Ikla core shows the prevalence of argillaceous marlstone containing nodules and interbeds of argillaceous limestone. The limestone interbeds are most common in the middle part (interval 256.2–262.4 m) of the Paramaja Member, which obviously corresponds to the Ninase Member developed in more northern sections. In the Ikla core, the Paramaja Member contains the stromatoporoids *Densastroma pexisum* (245.3 m), *Clathrodictyon linnarssoni* (246.5 and 261.5 m), and *Eostromatopora impexa* (246.5 m), indicating the shallowing-up trend of development.

Fig. 2. Lithological log and range chart of chitinozoan species of the Staicele drill core. Depths from the top of the drill hole are shown in metres. Lithological legend: 1, sandstone; 2, horizontal-bedded limestone; 3, nodular limestone; 4, wavy-bedded limestone; 5, limestone–marlstone intercalation; 6, wackestone; 7, packstone; 8, grainstone; 9, micritic limestone; 10, argillaceous limestone; 11, dolostone; 12, argillaceous dolostone; 13, marlstone; 14, highly dolomitic marlstone; 15, dolomitic marlstone; 16, argillaceous marlstone; 17, dolomitic argillaceous marlstone; 18, calcareous mudstone; 19, mudstone; 20, nodular dolomitic limestone; 21, marlstone with limestone nodules; 22, discontinuity surface (hardground); 23, metabentonite layer; 24, stromatolite with discontinuity surface; 25, red interval; 26, pyritized burrows; 27, fine bioclasts; 28, coarse bioclasts; 29, pyritized bioclasts; 30, pentamerid brachiopods; 31, pebbles; 32, pyritized pebbles; 33, oolites; 34, solution vugs; 35, pyrite mottles; 36, pyrite; 37, junction of formations; 38, junction of stages. Chitinozoan biozones: 1, *Conochitina alargada*; 2, *Eisenackitina dolioliformis*; 3, Interzone III; 4, *Conochitina mamilla*; 5, *Conochitina tuba*; 6, *Cingulochitina cingulata*; 7, *Eisenackitina lagena*; 8, *Conochitina pachycephala*. For regional stages see Fig. 4.

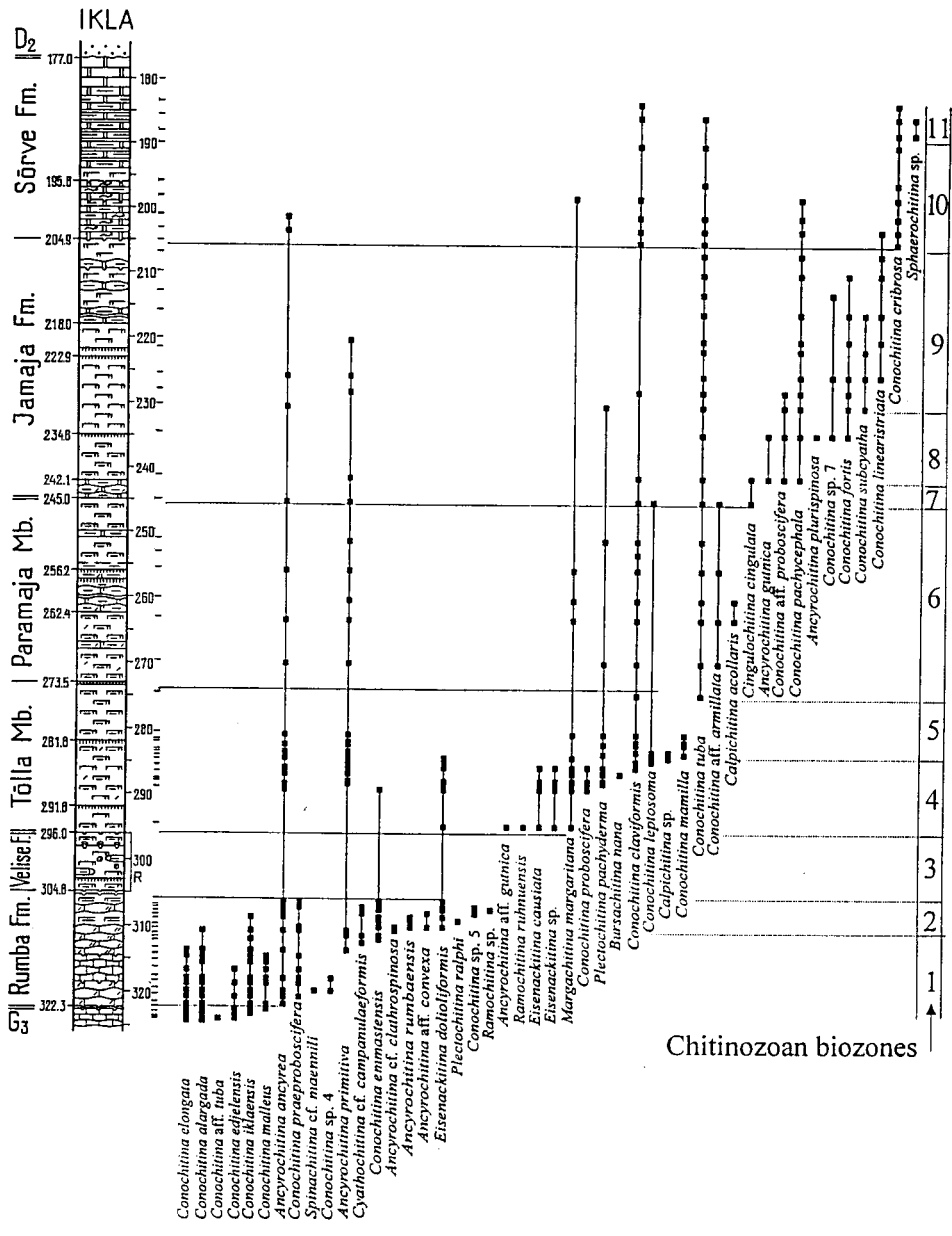


Fig. 3. Lithological log and range chart of chitinozoan species of the Ikla drill core. For lithological legend see Fig. 2. Chitinozoan biozones: 1, *Conochitina alargada*; 2, *Eisenackitina dolioliformis*; 3, Interzone III; 4, Interzone IV; 5, *Conochitina mamilla*; 6, *Conochitina tuba*; 7, *Cingulochitina cingulata* or *Eisenackitina lagena*; 8, *Conochitina pachycephala*; 9, *Conochitina subcyatha*; 10, *Conochitina cribrosa*; 11, *Sphaerochitina indecora*. For regional stages see Fig. 4.

Jaagarahu Stage

The Jaagarahu Stage is represented by the Jamaja Formation (below) and the Sõrve Formation (above). The rocks of the Sõrve Formation are preserved only in the Ikla core, but were eroded away completely in the vicinity of Staicele, in late- or post-Silurian time. The **Jamaja Formation** consists of variably dolomitized marlstones with a carbonate content increasing gradually upwards in the sequence. The marlstone is argillaceous in the lower part of the formation and calcareous in the upper part. It contains pyritized burrows and spots of dispersed pyrite, especially in the Staicele core. In the Ikla core, the uppermost part of the Jamaja Formation (interval 204.9–218.0 m) is characterized by cyclically alternating layers of calcareous marlstone and nodular argillaceous limestone, thus resembling the Riksu Formation developed to the northwest (H. Nestor et al. 2001). In the Ikla drill core a layer of nodular, dolomitized, argillaceous limestone occurs also at the base of the Jamaja Formation (interval 242.1–245.0 m). It contains skeletons of stromatoporoids which point to shallower-water conditions before or after a local hiatus, embracing a great part of the lowermost two Jaagarahu chitinozoan biozones in the Ikla core.

The uppermost portion of the incomplete Silurian sequence in the Ikla core (177.0–204.9 m) is represented by monotonous platy argillaceous dolomites. The dolomites are heavily bioturbated in their lower part (195.8–204.9 m), but very pure and thick-bedded in the uppermost part (177.0–181.0 m). In its stratigraphical position this dolomite unit corresponds to the **Sõrve Formation** but in terms of lithofacies it is quite different. The typical Sõrve Formation consists of nodular biomicritic limestones of the open shelf facies zone (according to H. Nestor & Einasto 1977). The platy dolomites in the Ikla core, however, are probably of restricted shelf genesis, resembling the Muhu Formation but lying much higher in the succession, at the level of a subregional hiatus established in most of the West Estonian sections (V. Nestor & H. Nestor 1991). Due to lithofacies differences these dolomite strata can be included in the Sõrve Formation only provisionally until their lateral extent is better studied.

The upper Silurian is missing in the considered sections. The rocks of the Jaagarahu Stage are unconformably overlain by the Lower Devonian Kemeru Stage (Väärssi et al. 1969).

SUCCESSION OF CHITINOZOANS

Adavere Stage

The chitinozoan range chart of the **Rumba Formation** in the Ikla core section has been published by V. Nestor (1994, fig. 13/3). It is used in the present paper alongside the new material and partly figured in Pls. I and II.

The lower part of the Rumba Formation of both studied cores is characterized by typical long-ranging middle Llandovery species *Conochitina edjelensis*

Taugourdeau, *C. elongata* Taugourdeau, *C. alargada* Cramer, *C. malleus* Van Grootel (*nomen nudum*), and *C. iklaensis* Nestor. The species *Conochitina praeproboscifera* Nestor, *Conochitina* sp. 4 (in V. Nestor 1994), and *Ancyrochitina primitiva* Eisenack appear in the lower–middle part of the Rumba Formation in Ikla (313–320 m) (Fig. 3) and in the middle–upper part of the same formation in Staicele (303–316 m) (Fig. 2). This assemblage is characteristic of the *Conochitina alargada* global Biozone (Verniers et al. 1995), which is usually correlated with the late Aeronian.

Explanation of Plate I

Fig. 1. *Conochitina alargada* Cramer 1967, 350-1, Staicele core, depth 321 m, Saarde Formation, Adavere Stage, ×140.

Fig. 2. *Conochitina iklaensis* Nestor 1980, 350-2, Ikla core, depth 320 m, Rumba Formation, Adavere Stage, ×115.

Fig. 3. *Conochitina malleus* (*nomen nudum*) Van Grootel 1990, 350-3, Staicele core, depth 320 m, Saarde Formation, ?Raikküla Stage, ×225.

Fig. 4. *Cyathochitina* cf. *campanulaeformis* (Eisenack 1931), 350-4, Ikla core, depth 310.6 m, Rumba Formation, Adavere Stage, ×170.

Fig. 5. *Conochitina praeproboscifera* Nestor 1994, 350-5, Staicele core, depth 298.2 m, Rumba Formation, Adavere Stage, ×220.

Fig. 6. *Ancyrochitina rumbaensis* Nestor 1994, 350-6, Ikla core, depth 307.8 m, Rumba Formation, Adavere Stage, ×155.

Fig. 7. *Ramochitina* sp., 350-7, Staicele core, depth 296 m, Rumba Formation, Adavere Stage, ×195.

Fig. 8. ?*Eisenackitina dolioliformis* Umnova 1976, 350-8, Staicele core, depth 295 m, Rumba Formation, Adavere Stage, ×125.

Fig. 9. *Eisenackitina causiata* Verniers 1999, 350-9, Ikla core, depth 294 m, Tõlla Member, Jaani Stage, ×400.

Fig. 10. *Eisenackitina* sp., 350-10, Ikla core, depth 294 m, Tõlla Member, Jaani Stage, ×255.

Fig. 11. *Margachitina margaritana* (Eisenack 1937), 350-11, Staicele core, depth 288.5 m, Tõlla Member, Jaani Stage, ×310.

Fig. 12. *Conochitina proboscifera* Eisenack 1937, 350-12, Ikla core, depth 287.5 m, Tõlla Member, Jaani Stage, ×70.

Fig. 13. *Plectochitina pachyderma* (Laufeld 1974), 350-13, Staicele core, depth 256.4 m, Jamaja Formation, Jaagarahu Stage, ×280.

Fig. 14. *Conochitina claviformis* Eisenack 1931, 350-14, Staicele core, depth 288.5 m, Tõlla Member, Jaani Stage, ×75.

Fig. 15. *Conochitina mamilla* Laufeld 1974, 350-15, Staicele core, depth 288.5 m, Tõlla Member, Jaani Stage, ×60.

Fig. 16. *Conochitina leptosoma* Laufeld 1974, 350-16, Staicele core, depth 280.3 m, Paramaja Member, Jaani Stage, ×75.

PLATE I

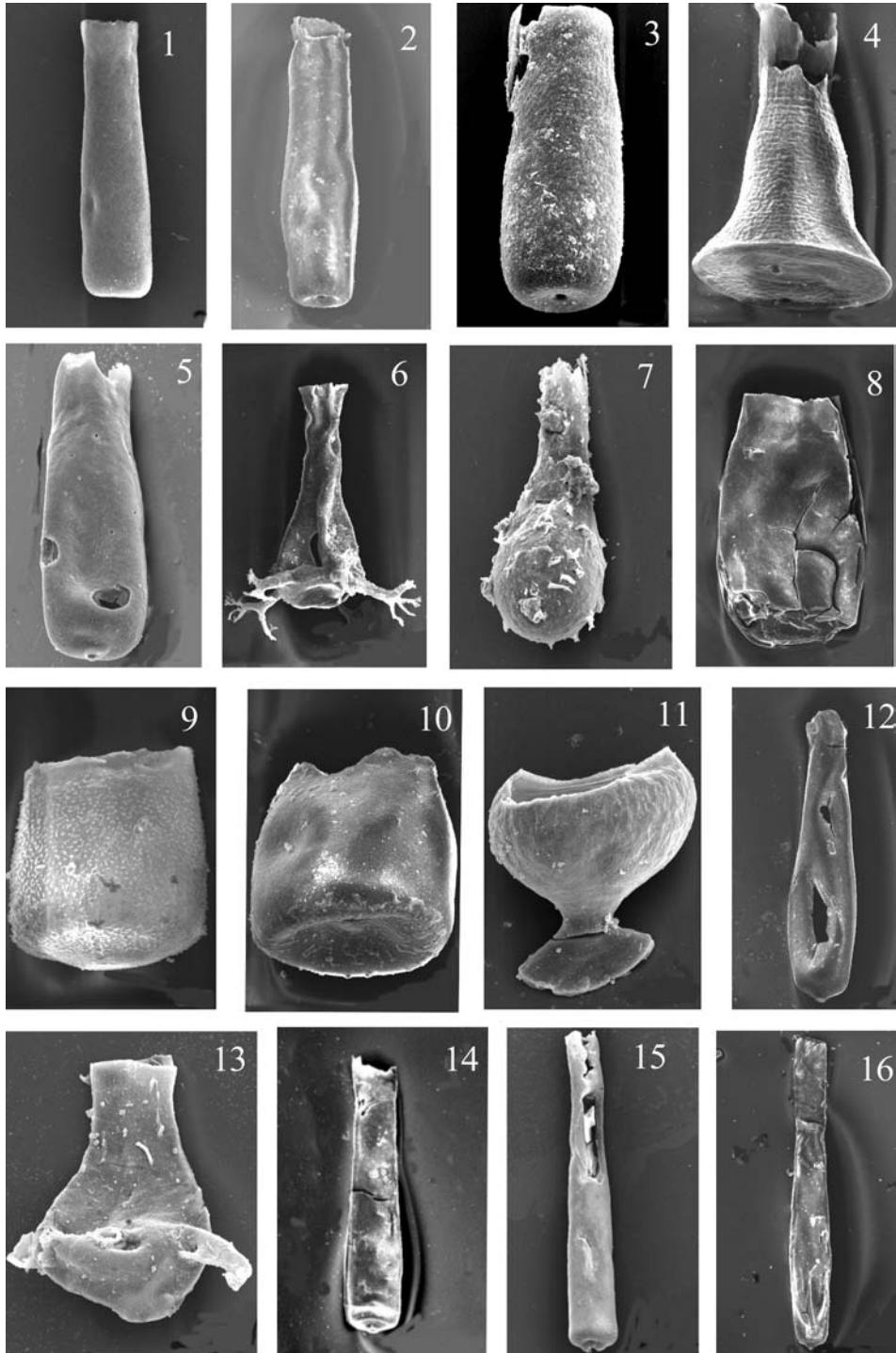
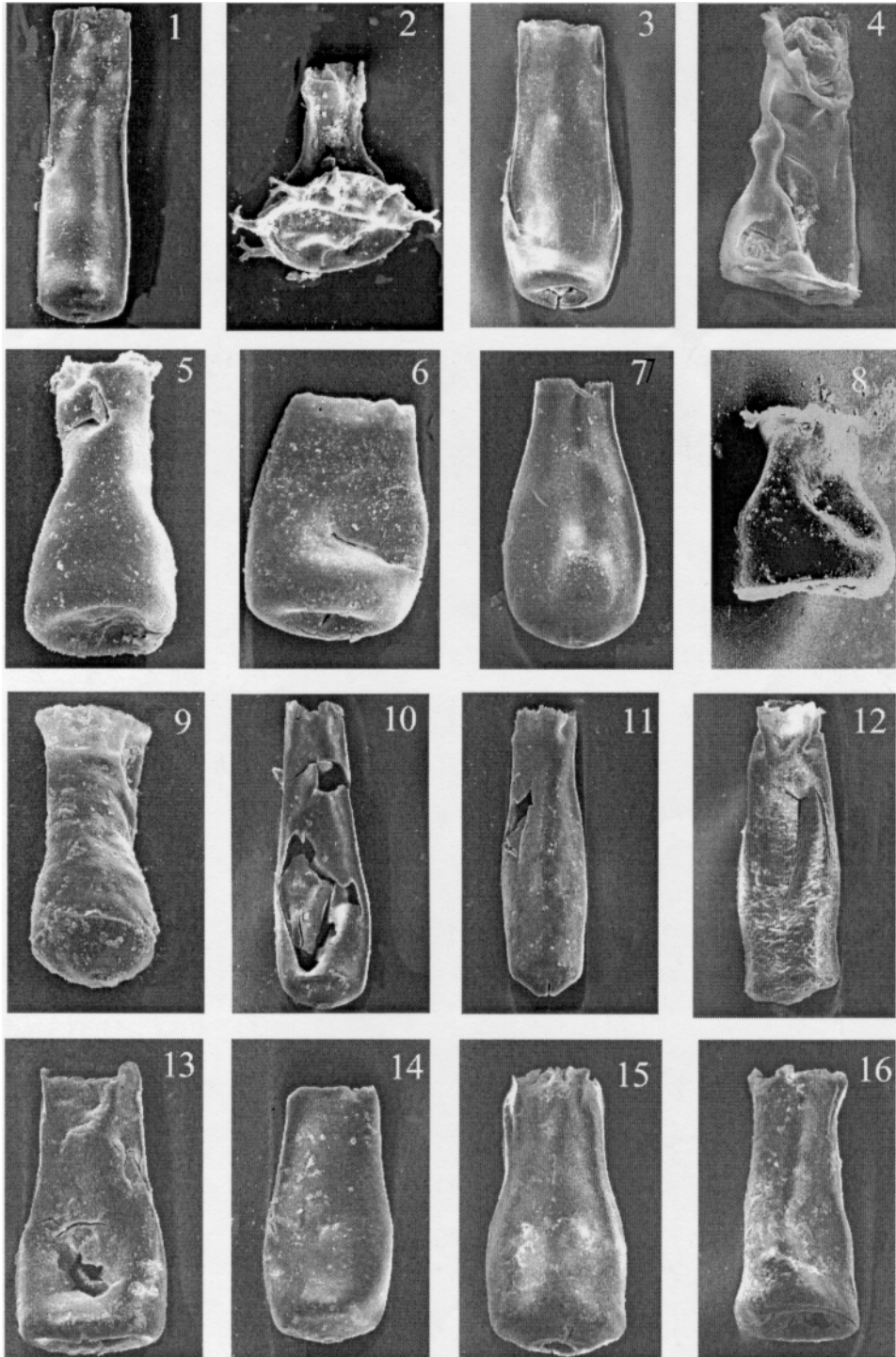


PLATE II



The zonal species of the next global biozone, *Eisenackitina dolioliformis* Umnova and *Conochitina emmastensis* Nestor, occur only in the upper part of the Rumba Formation, appearing first in the Ikla core at 309.3 and 311.2 m and in the Staicele core at 298.2 and 300 m, respectively. Also, short-ranging taxa such as *Ancyrochitina rumbaensis* Nestor and *Conochitina* sp. 5 (in V. Nestor 1994) are found in both cores and *Ancyrochitina* aff. *convexa* Nestor and *Plectochitina ralphi* Nestor in the upper portion of the Rumba Formation in Ikla. All long-ranging chitinozoan species, except *Ancyrochitina ancyrea* and *A. primitiva*, occurring in abundance throughout the Aeronian, disappear finally in the upper part of the Rumba Formation.

Explanation of Plate II

Fig. 1. *Conochitina tuba* Eisenack 1932, 350-17, Staicele core, depth 274.5 m, Paramaja Member, Jaani Stage, ×125.

Fig. 2. *Ancyrochitina paulaspina* Nestor 1974, 350-18, Staicele core, depth 270 m, Paramaja Member, Jaani Stage, ×265.

Fig. 3. *Conochitina* aff. *armillata* Taugourdeau & Jekhowski 1963, 350-19, Staicele core, depth 274.5 m, Paramaja Member, Jaani Stage, ×130.

Fig. 4. *Cingulochitina cingulata* (Eisenack 1937), 350-20, Staicele core, depth 266.2 m, Jamaja Formation, Jaagarahu Stage, ×300.

Fig. 5. *Conochitina* cf. *argillophila* Laufeld 1974, 350-21, Staicele core, depth 253.3 m, Jamaja Formation, Jaagarahu Stage, ×155.

Fig. 6. *Bursachitina* sp. 2, 350-22, Staicele core, depth 247.8 m, Jamaja Formation, Jaagarahu Stage, ×320.

Fig. 7. *Eisenackitina lagena* (Eisenack 1968), 350-23, Staicele core, depth 249 m, Jamaja Formation, Jaagarahu Stage, ×125.

Fig. 8. *Cingulochitina crassa* Nestor 1994, 350-24, Staicele core, depth 252.3 m, Jamaja Formation, Jaagarahu Stage, ×400.

Fig. 9. *Cingulochitina baltica* Nestor 1994, 350-25, Staicele core, depth 246.4 m, Jamaja Formation, Jaagarahu Stage, ×345.

Fig. 10. *Conochitina* aff. *proboscifera* Eisenack 1937, 350-26, Ikla core, depth 234 m, Jamaja Formation, Jaagarahu Stage, ×130.

Figs. 11, 12. *Conochitina pachycephala* Eisenack 1964, Ikla core, Jamaja Formation, Jaagarahu Stage: 11, 350-27, depth 234 m, ×110; 12, 350-28, depth 229.7 m, ×125.

Fig. 13. *Conochitina subcyatha* Nestor 1994, 350-29, Ikla core, depth 225 m, Jamaja Formation, Jaagarahu Stage, ×215.

Fig. 14. *Conochitina fortis* Nestor 1994, 350-30, Ikla core, depth 229.7 m, Jamaja Formation, Jaagarahu Stage, ×165.

Fig. 15. *Conochitina cribrosa* Nestor 1994, 350-31, Ikla core, depth 202.6 m, Sörve Formation, Jaagarahu Stage, ×175.

Fig. 16. *Conochitina linearistriata* Nestor 1994, 350-32, Ikla core, depth 209.5 m, Jamaja Formation, Jaagarahu Stage, ×180.

The highly reduced **Velise Formation** is represented by red beds, barren of chitinozoans. This lowermost part of the formation is distinguished as Interzone III (see V. Nestor 1994). The upper part of the formation corresponding to the *Angochitina longicollis* and *Conochitina proboscifera* local biozones is obviously lacking in both, cores due to a stratigraphical gap (see below).

Jaani Stage

In the considered sections, the **Riga Formation** of the Jaani Stage is represented by the Tõlla Member. *Margachitina margaritana* (Eisenack), *Ramochitina ruhnuensis* (Nestor), *Eisenackitina causiata* Verniers (= *Eisenackitina* sp. 1 in V. Nestor 1994) (from 294 m), and *Conochitina proboscifera* Eisenack (from 288.3 m) are present in the lowermost Tõlla beds of the Ikla core. However, *Angochitina longicollis* Eisenack was not found, which may indicate a stratigraphical gap in the Ikla core corresponding to the *Margachitina margaritana* local Biozone (V. Nestor 1994, p. 129). The occurrence of *C. proboscifera* and the absence of *A. longicollis* point to the presence of Interzone IV (V. Nestor 1994, p. 132) in the lowermost part of the Tõlla Member in the Ikla core. In the Staicele core, *E. causiata* and *C. proboscifera* are lacking and the Jaani succession begins with *Conochitina claviformis* Eisenack and *Conochitina mamilla* Laufeld (from 290.5 m) defining the base of the *C. mamilla* Biozone. The latter biozone is characteristic of the upper part of the Tõlla Member (V. Nestor 1994, p. 132). In the Ikla core, *C. mamilla* appears at 285 m. *Eisenackitina dolioliformis*, *E. causiata*, and *Conochitina proboscifera* disappear from the sequence at approximately the same level. Thus in the Staicele core the gap at the base of the Jaani Stage (and Wenlock) is bigger than in Ikla and embraces the *Margachitina margaritana* local Biozone and the succeeding Interzone IV (Fig. 4).

The **Jaani Formation** is represented by the Paramaja Member in both cores. Near the base of the member appear *Conochitina tuba* Eisenack (at 274 m in Ikla, at 280.3 m in Staicele) and *Calpichitina acollaris* (Eisenack) (at 262.5 m in Ikla, at 278.5 m in Staicele). The former species defines the *Conochitina tuba* Biozone (V. Nestor 1994, p. 134). The chitinozoan association of the Paramaja Member is dominated by *Conochitina claviformis* and *Margachitina margaritana*. The species *Plectochitina pachyderma* (Laufeld), *Conochitina* aff. *armillata* Taugourdeau, and *Ancyrochitina paulaspina* Nestor are less common. The last species occurs only in the Staicele core.

Jaagarahu Stage

The chitinozoan assemblage changes distinctly at the lower boundary of the **Jamaja Formation** in both cores. In the Staicele core, *Clathrochitina clathrata* Eisenack, *Ramochitina martinsoni* (Laufeld), and *Cingulochitina cingulata* (Eisenack) appear at a depth of 266.2 m. These species usually appear in the

SERIES	GLOBAL STAGE	REGIONAL STAGE	CHITINOZOAN BIOZONES V. Nestor 1994	Häädemeeste	Ikla	Staticele	Ristiküla	Ipiku		
WENLOCK	HOMERIAN	JAAGARAHU	<i>S. indecora</i>	117.5	177.0					
			<i>C. cribrosa</i>	S						
	SHEINWOODIAN	JAANI	<i>C. subcyatha</i>	130.9	204.9	235.0	119.8			
			<i>C. pachycephala</i>	J	J					
			<i>E. lagena</i>	J	J					
			<i>C. cingulata</i>	J	J					
	LANDOVERY	TELYCHIAN	ADAVERE	<i>C. clathrata</i>	121.9	245.0	266.0		241.6	
				<i>C. tuba</i>	P	P				
		?			<i>C. tuba</i>	195.0	273.5	280.5		
					<i>C. mamilla</i>	N	T			
Interzone IV					M	T				
<i>M. margaritana</i>					M	T				
			<i>C. proboscifera</i>	215.4	296.0					
			<i>A. longicollis</i>	V	V					
			Interzone III	V	V					
			<i>E. dolioformis</i>	R	R					
			<i>C. atargada</i>	218.7	304.8	292.6	180.5	260.8		
			<i>E. dolioformis</i>	218.7	304.8	295.0	180.5	260.8		
			<i>S. laevis</i>	218.7	304.8	295.0	180.5	260.8		
			<i>P. oblongus</i>	218.7	304.8	295.0	180.5	260.8		
			<i>C. atargada</i>	218.7	304.8	295.0	180.5	260.8		
				231.9	322.3	309.5	191.2	267.2		

Fig. 4. Correlation chart of the studied sections of southwestern mainland Estonia and northernmost Latvia with zonal succession of chitinozoans, stratigraphical gaps (hatched intervals), guide fossils, and specific lithological features. Depths of the junctions of stratigraphical units are given in metres from the top of the drill hole, species names refer to the appearance of zonal species above the base of the unit. Abbreviations of stratigraphical units: S, Sõrve Formation; J, Jamaja Formation; T, Tõlla Member; P, Paramaja Member; N, Ninase Member; M, Musijala Member; V, Velise Formation; R, Rumba Formation.

lowermost part of the Jamaja Formation corresponding to the *Cingulochitina cingulata* Biozone (V. Nestor 1994, p. 134). In the middle part of the Jamaja Formation, *Ancyrochitina plurispinosa* Nestor, *Conochitina* cf. *argillophila* Laufeld, *Cingulochitina crassa* Nestor, *Cingulochitina baltica* Nestor, and a zonal species *Eisenackitina lagena* (Eisenack) appear, the last one at a depth of 249 m, which marks the base of the *E. lagena* Biozone.

Conochitina pachycephala Eisenack as an index species and *Conochitina* aff. *proboscifera* Eisenack determine the next, *Conochitina pachycephala* Biozone (V. Nestor 1994, p. 136), occurring in the Staicele core in the uppermost part of the local Silurian sequence (244.5–246.5 m). This shows that temporal counterparts of the uppermost Jamaja (*Conochitina subcyatha* Biozone) and Sõrve formations are lacking in this core due to late or post-Silurian denudation.

In the Ikla core, an assemblage containing the zonal species *Conochitina pachycephala* and *C.* aff. *proboscifera* appears about 4 m above the base of the Jamaja Formation (at 240.8 m). *Cingulochitina cingulata*, which commonly ranges in the *Cingulochitina cingulata* and *Eisenackitina lagena* biozones, is present in this basal interval. This shows that the lower part of the Jamaja Formation, corresponding to the *Cingulochitina cingulata* and *Eisenackitina lagena* biozones, has a much reduced thickness in the Ikla core and may contain a local hiatus. However, no clear evidence was found to decide to which part of the *C. cingulata* and/or *E. lagena* biozones the supposed hiatus near the base of the Jamaja Formation exactly corresponds (Fig. 4). The upper part of the Jamaja Formation in the Ikla core (interval 206.5–234 m) contains the chitinozoan assemblage of the *Conochitina subcyatha* Biozone including *Conochitina fortis* Nestor, *C. subcyatha* Nestor, *C. linearistriata* Nestor, and *Conochitina* sp. 7 (V. Nestor 1994). This biozone is lacking in the less complete Staicele core (Fig. 4).

The typical assemblage of the **Sõrve Formation** has previously been established only in the Ohesaare and Ruhnu cores (V. Nestor 1994). In the Ikla core, the number of species is more restricted, but *Conochitina cribrosa* Nestor (from 204.5 m) and *Sphaerochitina* sp. (from 187.8 m) enable correlation of the lower part of the formation with the *Conochitina cribrosa* Biozone and the upper part very tentatively with the *Sphaerochitina indecora* Biozone.

In conclusion, most of the upper Aeronian to lower Homeric global chitinozoan biozones (Verniers et al. 1995), except the *Angochitina longicollis* Biozone, are represented in the Ikla and Staicele cores. The Aeronian *Conochitina alargada* global Biozone (= *Conochitina* cf. *protracta* Biozone and Interzone II in V. Nestor 1994) is probably represented by its uppermost part, and the *Eisenackitina dolioliformis* global Biozone by its lowermost part (= local *Conochitina emmastensis* Biozone in V. Nestor 1994). The red beds of the Telychian Velise Formation are barren of chitinozoans (local Interzone III). Therefore, the *Angochitina longicollis* global Biozone is lacking in both cores, as is the local *Conochitina proboscifera* Biozone (Fig. 4). In the East Baltic sections, the lower Sheinwoodian *Margachitina margaritana* global Biozone contains four

local biozones (V. Nestor 1994), from which Interzone IV, *Conochitina mamilla*, and *C. tuba* biozones are established in the Ikla and the last two also in the Staicele core. The lowermost, local *Margachitina margaritana* Biozone is not recognized in either the Ikla or the Staicele core. The upper Sheinwoodian *Cingulochitina cingulata* global Biozone, consisting of the local *Cingulochitina cingulata* and *Eisenackitina lagena* biozones, is fully present in Staicele but is represented only in a single sample in the Ikla core (Figs. 3 and 4). The lower Homeric *Conochitina pachycephala* global Biozone includes the *Conochitina subcyatha* and *C. cribrosa* local biozones, all of which are represented in the Ikla core. In the Staicele core only the *Conochitina pachycephala* local Biozone is recognized. A hiatus corresponds to the *C. subcyatha* and *C. cribrosa* local biozones in this core (Fig. 4). The topmost Wenlock global *Sphaerochitina lycoperdoides* Biozone in the East Baltic area may correspond to the *Sphaerochitina indecora* local Biozone (V. Nestor 1994, p. 140), but the badly preserved *Sphaerochitina* specimens in the uppermost part of the Ikla core did not enable precise identification of this zone.

STRATIGRAPHICAL GAPS

The study of the chitinozoan successions in the Ikla and Staicele drill cores revealed some extensive stratigraphical gaps (Fig. 4). The most prominent hiatus, between the Raikküla and Adavere regional stages, was recognized by V. Nestor (1976) when it became evident that in western Estonia the *Pentamerus oblongus* Beds or the Rumba Formation unconformably overlapped different subunits of the Raikküla Stage. In core sections on northwestern Saaremaa (Muratsi, Undva, Viki, and Eikla) the Rumba Formation lies on the Järva-Jaani Beds of the Nurmekund Formation which forms the lowermost part of the Raikküla Stage corresponding to the Rhuddanian (H. Nestor 1997, figs. 64 and 65). In a south-easterly direction the stratigraphical range of the hiatus decreases gradually. Thus the gap is almost unrecognizable in the Ikla and Staicele cores, where the nodular wackestones of the Rumba Formation lie upon the cryptocrystalline (micritic) limestones of the Staicele Member of the uppermost Saarde Formation. However, in the Staicele core the thickness of the both units is almost twice as great as in the Ikla core. This leads to a supposition that a stratigraphical gap may exist in the vicinity of Ikla, too, corresponding to the uppermost part of the Staicele Member and/or to the lowermost beds of the Rumba Formation in the Staicele core. There is no chitinozoan biostratigraphical evidence for this, however.

At present no direct evidence is available for exact timing of the late Raikküla sedimentation break in the most complete sections of southwestern mainland Estonia. Records of graptolites from the Staicele Member of the uppermost Saarde Formation are extremely rare. Kaljo (in Kaljo & Vingisaar 1969) recorded *Petalograptus minor* Elles and *Monograptus gemmatus* (Barrande) from the base of the Staicele Member in the Ikla core. Ulst (1973) has identified *Demi-*

rastrites ex gr. *triangulatus*, *Coronograptus gregarius* (Lapworth), *Orthograptus insectiformis* (Nicholson), and *Rastrites* aff. *longispinus* Perner from the lower part of the Staicele Member in the Staicele core, whereas the upper 45 m of the member contain only a few biserials. The absence of diagnostic species of the *Demirastrites convolutus* Biozone and presence of *Coronograptus gregarius* and *Demirastrites* ex gr. *triangulatus* suggest a pre-*convolutus* Zone age for the lowermost portion of the Staicele Member. Rubel (1977) has identified *Stricklandia lens progressa* Williams from the Staicele Member in the Ikla core (at 355.9 m). This subspecies occurs in the Llandovery type area of southern Wales in the upper Aeronian (Cocks 1989).

Only a few biserial graptolites have been recorded from the overlying Rumba Formation (Kaljo & Vingisaar 1969; H. Nestor 1972). The formation was provisionally correlated with the *Stimulograptus sedgwickii* graptolite Biozone (H. Nestor 1972) as it conformably underlies the Velise Formation containing *Spirograptus turriculatus* Biozone graptolites in the Ohesaare drill core. Recently, however, Loydell et al. (1998) established that only the upper part of the *Spirograptus turriculatus* Biozone (the *Torquigraptus proteus* Subzone) is represented in the Ohesaare core. Therefore it is possible that most of the Rumba Formation may actually correspond to the *Spirograptus guerichi* Biozone and to the lower part of the *Spirograptus turriculatus* Biozone. From the Rumba Formation of the Ikla core, in the interval of 309.3–321.1 m, Rubel (1977) has identified *Stricklandia lens ultima* (= *Stricklandia laevis* Sowerby), which in the Llandovery type area is diagnostic to the lowermost part of the Telychian Stage (Cocks 1989).

The above-stated observations, and first of all the distribution of the species of *Stricklandia*, allow us to conclude that the junction of the Saarde and Rumba formations, treated also as the boundary between the Raikküla and Adavere regional stages, lies close to the Aeronian–Telychian boundary in their type area in southern Wales. However, the junction of the *Conochitina alargada* and *Eisenackitina dolioliformis* global chitinozoan biozones, commonly treated as the Aeronian–Telychian boundary, lies higher in the Ikla and Staicele cores, within the Rumba Formation. This leaves the exact position of the Aeronian–Telychian boundary open. Loydell (1998) stated that a major eustatic fall in sea level occurred near the end of the Aeronian (during the *Stimulograptus sedgwickii* Chron). The hiatus below the Rumba Formation is presumably connected with this eustatic sea-level low-stand reflecting a glaciation in South America late in the Aeronian (Caputo 1998).

The present study of the chitinozoan successions in the Ikla and Staicele drill cores has revealed a stratigraphical gap near the junction of the Adavere and Jaani stages corresponding roughly to the Llandovery–Wenlock boundary (e.g. Kaljo 1962). In both cores, red to varicoloured marl- and mudstones, barren of fossils, of the Velise Formation (Adavere Stage) are overlain by grey dolomitic marlstones of the Tõlla Member of the Riga Formation (Jaani Stage). In the Staicele core, *Conochitina mamilla* and *C. claviformis*, indicative of the *Conochitina*

mamilla Biozone, occur at the base of the Tõlla Member (Fig. 4). Thus, the lowermost Wenlock *Margachitina margaritana* local Biozone and succeeding Interzone IV (V. Nestor 1994) are lacking in the Staicele core (Fig. 4). In the Ikla core, *C. claviformis* and *C. mamilla* appear about 10–12 m above the base of the Tõlla Member but *Angochitina longicollis*, which commonly ranges into the *Margachitina margaritana* Biozone from the upper Llandovery, was not found. Therefore the *Margachitina margaritana* local Biozone appears to be lacking in this core, and the Jaani Stage appears to begin at the level of Interzone IV (Fig. 4).

The underlying Velise Formation has a reduced thickness in both the Ikla and Staicele cores and shows some shallow-water lithological features (iron oolites, hardgrounds with pyritized pebbles, stromatolites). In some more complete, westerly drill cores of the Velise Formation (Kihnu, Seliste, Eikla, and Viki) red or varicoloured rocks occur only in the basal part of the Velise Formation. This leads to the conclusion that only the lower portion of the Velise Formation is preserved in the Ikla and Staicele cores, while the upper part was not deposited or was eroded away during the end-Telychian to earliest Wenlock break in sedimentation.

Comparison of the Ikla and Staicele cores shows that the gap between the Adavere and Jaani Stages increases towards the east (Fig. 4). In the Ristiküla and Ipiku drill cores (Fig. 1), studied during the geological mapping of the territory (Väärsi et al. 1969), the Velise Formation is totally lacking, and marlstones of the Tõlla Member, with the zonal graptolite *Monograptus riccartonensis*, directly overlie the nodular wackestones of the Rumba Formation (Fig. 4). Thus, in these sections the hiatus embraces the interval from the *Spirograptus turriculatus* to the *Monograptus firmus* graptolite biozones.

The stratigraphical range of the hiatus decreases to the west and northwest. In the Häädemeeste core the Velise Formation has still a reduced thickness and is represented by red beds. The Jaani Stage begins with the *Margachitina margaritana* local chitinozoan Biozone defined by the co-occurrence of *M. margaritana* and *Angochitina longicollis* (V. Nestor 1994, fig. 12/3). Westwards and northwestwards the thickness of the Velise Formation increases markedly and no direct evidence of the hiatus has been found. However, the top-most Llandovery *Cyrtograptus insectus* graptolite Biozone has not been established anywhere in the East Baltic area, not even in the most complete graptolitic sections of the Baltic Syneclise (Paškevičius 1998). The basal Wenlock *Cyrtograptus centrifugus* Biozone has been recognized only in the Kurtuvėnai-161 borehole in western Lithuania (Paškevičius 1998), where it overlies the *Cyrtograptus lapworthi* Biozone. In the Ohesaare core in southwestern Saaremaa, graptolitic mudstones of early *Cyrtograptus lapworthi* Biozone age are overlain by 7.7 m thick non-graptolitic greenish-grey mudstones, succeeded by dark mudstones with a *Cyrtograptus purchisoni* Biozone fauna (Loydell et al. 1998). These data show a general sea-level low-stand close to the Llandovery–Wenlock boundary which

in places may have resulted in a break in sedimentation. In several drill cores (Kihnu, Varbla, Lihula, and Suuremõisa) E. Kala and H. Perens (pers. comm.) discovered scattered lime oolites in the basal layer of the Jaani Formation. These were probably brought in from adjacent shallow-water shoal areas at the beginning of the early Wenlock transgression that reached its maximum in the East Baltic region by the *Monograptus riccartonensis* Chron (Bassett et al. 1989, p. 161; Paškevičius 1998, p. 122). At that time sedimentation recommenced in the vicinity of Ipiku and Ristiküla, and somewhat later, during the *Conochitina mamilla* Chron, in Staicele (Fig. 4).

To a certain extent, the hiatus in southwestern mainland Estonia and northernmost Latvia coincides with the Ireviken conodont extinction event (Jeppsson & Männik 1993; Jeppsson et al. 1995) and also with a prominent positive carbon isotope excursion in the lowermost Wenlock of the Ohesaare, Ruhnu, Ventspils, and Priekule drill cores (Kaljo et al. 1997, 1998; Kaljo & Martma 2000). The data above support the opinion of Loydell (1998) of a major sea-level drop just prior to the Llandovery–Wenlock transition, which was probably related to a glaciation episode in the Amazon Basin of Brazil at the end of the accumulation of the Nhamundá Formation (Caputo 1998).

One more stratigraphical gap was established in the Ikla core, at the base of the Jamaja Formation, where a great part of the lower Jaagarahu *Cingulochitina cingulata* and *Eisenackitina lagenae* local chitinozoan biozones is lacking (Fig. 4). The hiatus has not been detected in the adjacent Staicele and Häädemeeste drill cores and has probably only local significance. In southwestern mainland Estonia and northernmost Latvia, the Silurian sequence ends at different levels in the upper Wenlock, which is unconformably overlain by the Lower Devonian Kemeru Stage.

CONCLUSIONS

Three stratigraphical gaps were established at the boundaries of regional stages in the lower Silurian sections of southwestern mainland Estonia and northernmost Latvia. The hiatuses at the junctions of the Raikküla–Adavere and Adavere–Jaani stages have a considerable areal extent and stratigraphical range, particularly the former. They are probably connected with the global glacio-eustatic sea-level falls near the end of the Aeronian and Telychian (Loydell 1998) corresponding to glacial episodes in the Amazon Basin of Brazil (Caputo 1998). The third gap, near the base of the Jaagarahu Stage (Jamaja Formation), has been recognized only in the Ikla core and has local significance.

The Rumba Formation of the Adavere Stage, hitherto indirectly correlated with the *Stimulograptus sedgwickii* graptolite Zone and assigned to the Aeronian, belongs, at least partly, to the Telychian based on the presence of the brachiopod *Stricklandia laevis* in the Ikla core, and *Eisenackitina dolioliformis* Zone chitinozoans in the upper part of the formation.

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REFERENCES

- Aaloe, A. 1970. Jaani Stage. Jaagarahu Stage. In *Silur Estonii* (Kaljo, D., ed.), pp. 243–264. Valgus, Tallinn (in Russian).
- Bassett, M. G., Kaljo, D. & Teller, L. 1989. The Baltic region. In *A Global Standard for the Silurian System* (Holland, C. H. & Bassett, M. G., eds.). *Nat. Mus. Wales Geol. Ser.*, **9**, 158–170.
- Caputo, M. V. 1998. Ordovician–Silurian glaciations and global sea-level changes. In *Silurian Cycles. Linkages of Dynamic Stratigraphy with Atmospheric, Oceanic, and Tectonic Changes* (Landing, E. & Johnson, M. E., eds.). *New York State Mus. Bull.*, **491**, 15–25.
- Cocks, L. R. M. 1989. The Llandovery Series in the Llandovery area. In *A Global Standard for the Silurian System* (Holland, C. H. & Bassett, M. G., eds.). *Nat. Mus. Wales Geol. Ser.*, **9**, 36–50.
- Einasto, R. 1984. Main stages of development and facies models of the East Baltic Silurian peri-continental basin. In *Theory and Practice of Ecostratigraphy* (Kaljo, D. & Klaamann, E., eds.), pp. 37–54. Valgus, Tallinn (in Russian).
- Einasto, R., Nestor, H., Kala, E. & Kajak, K. 1972. Correlation of the Upper Llandoveryan sections in West Estonia. *ENSV TA Toim. Keemia Geol.*, **21**, 333–343 (in Russian).
- Gailite, L. K., Ulst, R. J. & Jakovleva, V. I. 1987. *Stratopicheskie i tipovye razrezy Latvii*. Zinatne, Riga (in Russian).
- Jeppsson, L. & Männik, P. 1993. High resolution correlations between Gotland and Estonia near the base of the Wenlock. *Terra Nova*, **5**, 348–358.
- Jeppsson, L., Aldridge, R. J. & Dorning, K. J. 1995. Wenlock (Silurian) oceanic episodes and events. *J. Geol. Soc. London*, **152**, 487–498.
- Kaljo, D. 1962. On the boundary of the Llandoveryan and Wenlockian in the East Baltic. *ENSV TA Geol. Inst. Uurimused*, **10**, 97–113 (in Russian).
- Kaljo, D. & Martma, T. 2000. Carbon isotopic composition of Llandovery rocks (East Baltic Silurian) with environmental interpretation. *Proc. Estonian Acad. Sci. Geol.*, **49**, 267–283.
- Kaljo, D. & Vingisaar, P. 1969. On the sequence of the Raikküla Stage in southernmost Estonia. *ENSV TA Toim. Keemia Geol.*, **18**, 270–277 (in Russian).
- Kaljo, D., Kiipli, T. & Martma, T. 1997. Carbon isotope event markers through the Wenlock–Přidoli sequence at Ohesaare (Estonia) and Priekule (Latvia). *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, **132**, 211–223.
- Kaljo, D., Kiipli, T. & Martma, T. 1998. Correlation of carbon isotope events and environmental cyclicity in the East Baltic Silurian. In *Silurian Cycles. Linkages of Dynamic Stratigraphy with Atmospheric, Oceanic, and Tectonic Changes* (Landing, E. & Johnson, M. E., eds.). *New York State Mus. Bull.*, **491**, 297–312.
- Loydell, D. K. 1998. Early Silurian sea-level changes. *Geol. Mag.*, **135**, 447–471.
- Loydell, D. K., Kaljo, D. & Männik, P. 1998. Integrated biostratigraphy of the lower Silurian of the Ohesaare core, Saaremaa, Estonia. *Geol. Mag.*, **135**, 769–783.
- Nestor, H. 1972. On the stratigraphic range of the beds with *Pentamerus oblongus* and on the nature of the late Llandoveryan transgression in North Europe. *ENSV TA Toim. Keemia Geol.*, **21**, 344–350 (in Russian).

- Nestor, H. 1997. Silurian. In *Geology and Mineral Resources of Estonia* (Raukas, A. & Teedumäe, A., eds.), pp. 89–106. Estonian Acad. Publ., Tallinn.
- Nestor, H. & Einasto, R. 1977. Facies-sedimentary model of the Silurian Paleobaltic pericontinental basin. In *Facies and Fauna of the Baltic Silurian* (Kaljo, D., ed.), pp. 89–121. Acad. Sci. Estonian SSR, Tallinn (in Russian).
- Nestor, H. & Einasto, R. 1982. Application of shelf and slope concepts to the Silurian Baltic basin. In *Ecostratigraphy of the East Baltic Silurian* (Kaljo, D. & Klaamann, E., eds.), pp. 17–24. Acad. Sci. Estonian SSR, Tallinn.
- Nestor, H., Einasto, R., Nestor, V., Märss, T. & Viira, V. 2001. Description of the type section, cyclicity, and correlation of the Riksu Formation (Wenlock, Estonia). *Proc. Estonian Acad. Sci. Geol.*, **50**, 149–173.
- Nestor, V. 1976. A microplankton correlation of boring sections of the Raikküla Stage, Estonia. *ENSV TA Toim. Keemia Geol.*, **25**, 319–324 (in Russian).
- Nestor, V. 1994. Early Silurian chitinozoans of Estonia and North Latvia. *Academia*, **4**. Estonian Acad. Publ., Tallinn.
- Nestor, V. & Nestor, H. 1991. Dating of the Wenlock carbonate sequences in Estonia and stratigraphic breaks. *Proc. Estonian Acad. Sci. Geol.*, **40**, 50–60.
- Paškevičius, J. 1998. *The Geology of the Baltic Republics*. Vilnius Univ., Geol. Surv. Lithuania, Vilnius.
- Rubel, M. 1977. Evolution of the genus *Stricklandia* (Pentamerida, brach.) in the Llandovery of Estonia. In *Facies and Fauna of the Baltic Silurian* (Kaljo, D., ed.), pp. 193–212. Acad. Sci. Estonian SSR, Tallinn (in Russian).
- Ulst, R. 1973. Llandoveryan deposits of northern Latvia and their correlation. In *Problemy regional'noj geologii Pribaltiki i Belorussii* (Kuršs, V. M., ed.), pp. 71–85. Zinatne, Riga.
- Väärsi, A. R., Kajak, K. F., Kajak, H. V., Kirs, E. E. & Liivrand, H. J. 1969. Report of the South-Estonian Geological Squad on complex geological-hydrogeological mapping on the scale 1 : 200 000 of the southwestern part of Estonia (card O-35-XIII) for 1966–1968. Geol. Surv. Estonian SSR, Keila (unpublished report in Russian).
- Verniers, J., Nestor, V., Paris, F., Dufka, P., Sutherland, S. & van Grootel, G. 1995. A global Chitinozoa biozonation for the Silurian. *Geol. Mag.*, **132**, 651–666.

Ülem-Llandovery–Kesk-Wenlocki (Silur) litostratigraafia ja kitinosaade biostratigraafia Edela-Eestis ja Põhja-Lätis

Heldur Nestor ja Viiu Nestor

On esitatud andmed Adavere, Jaani ja Jaagarahu lademe litostratigraafilise liigestuse ning kitinosaade leviku kohta Ikla ja Staicele puursüdamikus. Adavere lademe Rumba kihistu on vähemalt osaliselt Telychi vanusega. Redutseerunud paksusega Velise kihistu punavärvilised merglid esindavad vaid kihistu alumist osa. Settelüngad Raikküla–Adavere ja Adavere–Jaani lademe vahel on märkimisväärselt pindalalise leviku ja stratigraafilise ulatusega ning peegeldavad tõenäoliselt glatsio-eustaatilisi merepinna langusi, mis vastavad mandrijäätumistele Lõuna-Ameerikas Aeroni ja Telychi lõpul. Jaagarahu lademe alumisele osale vastav lünk Ikla puuraugus on seevastu lokaalse tähendusega.

Литостратиграфия и распространение хитиной в верхнем лландовери–среднем венлоке (силур) Юго-Западной Эстонии и Северной Латвии

Хелдур Нестор и Вийу Нестор

Приведены данные о литостратиграфическом расчленении и распространении хитиной в адавереском, яанском и яагарахуском горизонтах в скважинах Икла и Стайцэле. Румбаская свита адавереского горизонта относится, во всяком случае частично, к подъярису Телич. Маломощные красноцветные мергели велизской свиты представлены лишь в нижней части свиты. Стратиграфические перерывы между райккюласким и адавереским, а также между адавереским и яанским горизонтами имеют широкое площадное распространение и заметный стратиграфический диапазон. Они, вероятно, отражают гляцио-эвстатические понижения уровня мирового океана, связанные с оледенениями в Южной Америке. Стратиграфический перерыв в основании яагарахуского горизонта в скважине Икла имеет локальное значение.