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## COMPARISON OF TWO METHODS FOR THE CALCULATION OF THE ATMOSPHERIC INTEGRAL TRANSPARENCY COEFFICIENT

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Abstract. Accuracy of the transformation (reduction) of the atmospheric integral transparency coefficient from the actual air mass m to air mass m = 2 (solar elevation  $h = 30^{\circ}$ ) was studied using two different methods – the method proposed by Evnevich–Savikovskij, and the method proposed by Mürk–Ohvril. The table of average values of direct solar radiation versus solar elevation angle and rate of atmospheric turbidity, compiled by Sivkov and updated by Evnevich–Savikovskij, was considered as the basic one. The results show that both methods give almost the same uncertainty (2%) and they may be recommended for practical calculations of the atmospheric integral transparency coefficient.

Key words: direct solar radiation, atmospheric integral transparency coefficient, Forbes' effect.

#### **INTRODUCTION**

The Atmospheric Integral Transparency Coefficient (AITC),  $p_m$ , is easily calculated, according to the well-known Bouguer law, from the measured integral (broadband) direct solar irradiance on a plane perpendicular to the beam,  $S_m$ :

$$p_m = \left(\frac{S_m}{S_0}\right)^{\frac{1}{m}}.$$
(1)

Here *m* is the number of (relative optical) air masses in the direction of the sun and  $S_o$  is the extraterrestrial irradiance determined by the earth-sun distance in astronomical units  $\rho$  and the solar constant  $S_o^* = 1.367 \text{ kW m}^{-2}$ :

$$S_{\rm o} = S_{\rm o}^* \frac{1}{\rho^2} \,. \tag{2}$$

The coefficient  $p_m$  is one of the simplest characteristics of the turbidity of the atmosphere. This is not a new parameter. Numerous actinometrists, mainly in the former USSR (FSU), have used it since the 1920s (Kondratyev, 1969). A large amount of measured and archived data on direct solar radiation in various meteorological stations over the world would be a rich database for the retrieval of long-term time series of AITC and the creation of an extremely useful climatological material of the local and global climate changes. Unfortunately, outside the FSU this parameter was almost unknown. Even in the FSU the number of meteorological stations where direct solar irradiance,  $S_m$ , was in the daily program of measurements was considerably reduced in the 1960s.

Mainly two kinds of difficulties are encountered in using  $S_m$  and  $p_m$ :

(1) high quality pyrheliometers used in measuring direct solar radiation are expensive but vulnerable instruments, their application in automatized measurements always contains risk of damage by precipitation; unfortunately cheap and more robust Actinometers AT50 (developed at the Main Geophysical Observatory in Leningrad), considered to provide acceptable accuracy, are hardly available now;

(2) because of the Forbes' effect (caused by the selective spectral attenuation of direct solar radiation in the atmosphere) the AITC depends on solar elevation even in the case of stationary and azimuthally homogeneous atmosphere.

The problem of vulnerability of pyrheliometers needs to be solved by the development of instruments of more reliable design. Until new instruments are available, the solutions might be: (1) manual use of pyrheliometers, or (2) measurement of  $S_m$  by pyranometers as the difference of global and diffuse solar radiation.

To eliminate the Forbes' effect, the generally accepted practice is to reduce the AITC  $p_m$  from the actual air mass m to  $p_2$  corresponding to the air mass m = 2 (solar elevation angle  $h = 30^\circ$ ). The complicated problem of reducing the Forbes' effect was in principle solved in the 1960s by Sergei Sivkov (who compiled tables for the transformation of AITC) and Herman Mürk (who created a nomogram for the same purpose).

Unfortunately both methods were quite clumsy, needing a certain amount of "manual work", which seems to be one of the reasons for their limited use. However, both methods have been considerably upgraded to simple analytical forms. Sivkov's method was developed to be more user-friendly by Evnevich & Savikovskij (Евневич & Савиковский, 1989) and Mürk's method by Mürk & Ohvril (Мюрк & Охвриль, 1988, 1990).

The method of Evnevich & Savikovskij allows one to calculate  $p_2$  directly from measured values of  $S_m$  using one of the two formulas presented below. The *first formula* (which we further call *model ES-1*):

$$p_2 = 0.978 \left(\frac{S_m}{1.307}\right)^{\frac{\sin h + 0.15}{1.3}},\tag{3}$$

or the second formula (model ES-2):

$$p_2 = \left(\frac{S_m}{1.367}\right)^{\frac{\sin h + 0.205}{1.41}}.$$
(4)

Mürk & Ohvril proposed a general formula to go from the transparency coefficient  $p_m$  to  $p_i$  (model MO-1):

$$p_{i} = p_{m} \left(\frac{i}{m}\right)^{\frac{\log p_{m} + 0.009}{\log m - 1.848}},$$
(5)

which in the most important particular case i = 2 ( $h = 30^{\circ}$ ) takes the form:

$$p_2 = p_m \left(\frac{2}{m}\right)^{\frac{\log p_m + 0.009}{\log m - 1.848}}.$$
(6)

The aim of the present paper is to compare the accuracy of these different methods for the calculation of AITC  $p_2$ .

# STANDARD VALUES OF DIRECT SOLAR IRRADIANCE

The greatest work in the generalization of measurements of direct solar irradiance,  $S_m$ , was done by Sivkov (Сивков, 1965). Using databases from eight meteorological stations on the territory of the USSR containing more than 13 000 measurements at different solar elevations  $(7^{\circ} \le h \le 42^{\circ} \text{ or } 8 \ge m \ge 1.5)$  and a wide range of atmospheric turbidity, he compiled a table of mean values of direct solar irradiance,  $S_m$ . He found the values for  $S_m$  at solar elevations  $h > 42^{\circ}$  using mathematical extrapolation.

In connection with the introduction of new radiation units, a new solar constant, and changes in the pyrheliometric scale, Evnevich published a new version of tables for  $S_m$  (Евневич, 1986) and then, together with Savikovskij, the newest version (Евневич & Савиковский, 1989), which we consider as standard (see Table 1). This table presents mean values of  $S_m$  corresponding to 10 different situations of atmospheric turbidity – from very low transparency ( $p_2 = 0.410$ ) to the ideal atmosphere ( $p_2 = 0.905$ ). All values are original except the value  $S_m = 1.082$  in the second row from bottom ( $p_2 = 0.872$ ,  $h = 40^\circ$ ), written in bold. In the original this value was 1.032, which is evidently a typographical error.

Table 1

Transparency,		Solar elevation angle, h									
p	2	10	20	30	40	50	60	70	80	80         90           0.475         0.486           0.554         0.564           0.691         0.704	
Very	0.410	0.035	0.133	0.230	0.316	0.377	0.426	0.456	0.475	0.486	
low	0.469	0.070	0.188	0.300	0.391	0.456	0.506	0.537	0.554	0.564	
Low	0.567	0.147	0.321	0.440	0.530	0.600	0.649	0.677	0.691	0.704	
	0.623	0.216	0.416	0.530	0.624	0.694	0.740	0.755	0.768	0.775	
Normal	0.700	0.346	0.551	0.670	0.756	0.810	0.846	0.866	0.879	0.886	
	0.736	0.412	0.621	0.740	0.817	0.865	0.900	0.914	0.928	0.935	
	0.770	0.489	0.698	0.810	0.879	0.921	0.956	0.970	0.984	0.986	
High	0.811	0.579	0.796	0.900	0.956	0.991	1.019	1.033	1.040	1.047	
	0.872	0.761	0.956	1.040	1.082	1.103	1.124	1.138	1.145	1.152	
Ideal	0.905	0.887	1.050	1.120	1.156	1.176	1.190	1.204	1.212	1.219	

Standard values of direct solar irradiance,  $S_m$ , in kW/m<sup>2</sup> according Evnevich & Savikovskij (Евневич & Савиковский, 1989)

In order to visualise the variation of standard values of  $S_m$  as functions of solar elevation angle, h, we have inserted Fig. 1, which contains 10 different plots according to Table 1.





### Comparison of methods

In order to compare the accuracy of the use of the three models discussed (ES-1, ES-2, and MO-1) we have proceeded from Table 1 and calculated for each standard  $S_m$  a standard value of AITC  $p_m$  according to formula (1). Then, proceeding from these models and the given AITC  $p_2$ , we calculated AITC  $p_m$  in three different ways:

(1) according to ES-1; revealing  $S_m$  from (3) and applying (1), we easily obtain

$$p_m = \left[ 0.9561 \left( \frac{p_2}{0.978} \right)^{\frac{1.3}{\sin h + 0.15}} \right]^{\frac{1}{m}},\tag{7}$$

(2) according to ES-2; revealing  $S_m$  from (4) and using again (1), we obtain

$$p_m = (p_2)^{\frac{1.41}{(\sin h + 0.205)m}},$$
(8)

(3) according to MO-1; from the general formula (5) follows a special one

$$p_m = p_2 \left(\frac{m}{2}\right)^{-\frac{\log p_2 + 0.009}{1.547}}.$$
(9)

The accuracy of the values of  $p_m$ , found by each model (ES-1, ES-2, MO-1) with regard to standard values,  $p_m$  (standard), was estimated by calculating the relative errors,  $\delta$ :

$$5 = \frac{p_m - p_m(\text{standard})}{p_m(\text{standard})} \cdot 100\% .$$
(10)

For three selected cases of atmospheric turbidity ( $p_2 = 0.410, 0.700, 0.872$ ) the results of the calculation of  $p_m$  and  $\delta$  are presented in Table 2.

To economize the space of this paper, the remaining seven cases of turbidity given in Table 1 are not presented. The greatest errors caused by each considered model occur in the case of very low transparency, when  $p_2 = 0.410$ .

The greatest of all errors, equal to 6.4%, is in the use of model ES-2 at solar elevation  $h = 10^{\circ}$ . The maximum error caused by model ES-1 is only 2.31% ( $h = 10^{\circ}$ ) and by model MO-1, 2.65% ( $h = 90^{\circ}$ ). In Table 2 the maximum errors in the case of  $p_2 = 0.410$  are written in bold.

It is important to stress that for solar elevations  $20^{\circ} \le h \le 80^{\circ}$ , all three models secure accuracy of reduction of AITC  $p_2$  to  $p_m$  with an error smaller than 2.5% in regard to standard values of  $p_m$ .

Comparison of method

Table 2

	old St	andard ai	nd modell	ed values	of AITC p	<sub>m</sub> with re	elative err	ors o	
h	m	S <sub>m</sub> (stand- ard)	$p_m$ (stand-ard)	ES-1		ES-2		MO-1	
				$p_m$	δ, %	$p_m$	δ, %	<i>p</i> <sub>m</sub>	δ, %
. We									
			Very	low transpa	arency, $p_2$	= 410			
10	5.60	0.035	0.520	0.532	2.31	0.553	6.40	0.527	1.47
20	2.90	0.133	0.448	0.446	-0.44	0.453	1.17	0.449	0.27
30	2.00	0.230	0.410	0.410	0.00	0.410	0.00	0.410	0.00
10	1.556	0.316	0.390	0.389	-0.41	0.386	-1.06	0.386	-1.16
50	1.305	0.377	0.373	0.375	0.69	0.371	-0.45	0.369	-0.89
50	1.155	0.426	0.364	0.367	0.72	0.362	-0.66	0.358	-1.62
70	1.064	0.456	0.356	0.362	1.47	0.356	0.00	0.351	-1.40
30	1.015	0.475	0.353	0.359	1.59	0.353	0.02	0.347	-1.59
00	1.00	0.486	0.356	0.358	0.64	0.352	-0.99	0.346	-2.65
			Norm	nal transpa	rency, $p_2$	= 0.700			
0	5.60	0.346	0.782	0.780	-0.26	0.789	0.90	0.771	-1.41
20	2.90	0.551	0.731	0.726	-0.68	0.728	-0.41	0.725	-0.82
30	2.00	0.670	0.700	0.700	0.00	0.700	0.00	0.700	0.00
40	1.556	0.756	0.683	0.683	0.00	0.683	0.00	0.684	0.15
50	1.305	0.810	0.670	0.672	0.30	0.672	0.30	0.672	0.30
50	1.155	0.846	0.660	0.664	0.61	0.666	0.91	0.665	0.76
70	1.064	0.866	0.651	0.659	1.23	0.662	1.69	0.660	1.38
30	1.015	0.879	0.647	0.656	1.39	0.659	1.85	0.657	1.55
00	1.00	0.886	0.648	0.655	1.08	0.659	1.70	0.656	1.23
			Very l	nigh transp	arency, p	2 = 0.872			
0	5.60	0.761	0.901	0.914	1.44	0.913	1.37	0.902	0.12
20	2.90	0.956	0.884	0.887	0.32	0.885	0.11	0.883	-0.15
30	2.00	1.040	0.872	0.872	0.00	0.872	0.00	0.872	0.00
10	1.556	1.082	0.860	0.861	0.04	0.864	0.41	0.865	0.51
0	1.305	1.103	0.848	0.853	0.52	0.859	1.25	0.860	1.36
50	1.155	1.124	0.844	0.847	0.34	0.855	1.29	0.857	1.47
70	1.064	01.138	0.842	0.843	0.14	0.853	1.34	0.854	1.49
30	1.015	1.145	0.840	0.840	0.00	0.852	1.45	0.853	1.56
0	1.00	1.152	0.843	0.840	-0.36	0.852	1.10	0.852	1.16

Figure 2 contains eight plots of reduced coefficients of  $p_m$  calculated by the three models considered, and plots of standard values of  $p_m$ (continuous line). It is obvious from the physical content of the Forbes' effect that the course of  $p_m$  as a function of diminishing solar elevation angles, h (or increasing optical masses, m), should form an increasing sequence

$$p_m (h = 90^\circ) < p_m (h = 80^\circ) < p_m (h = 70^\circ) < \dots < p_m (h = 10^\circ)$$
. (11)

All models used fulfil this demand except standard values of  $p_m$  at high solar elevations ( $h > 60^\circ$ ).



Fig. 2. Values of AITC  $p_m$  as functions of diminishing solar elevation angle, h.

For example, in the case of very low transparency ( $p_2 = 0.410$ ), standard values of  $p_m$  are

$$p_m (h = 90^\circ) = 0.356 > p_m (h = 80^\circ) = 0.353,$$
 (12)

and in the case of the ideal atmosphere ( $p_2 = 0.905$ ) even:

$$p_m (h = 90^\circ) > p_m (h = 60^\circ)$$
. (13)

Neither inequality (12) nor (13) is physically based.

Irregularities in the courses of standard values of  $p_m$  may be corrected by correcting standard values of direct solar irradiance,  $S_m$ , in Table 1. As mentioned above, Sivkov (Сивков, 1965) found standard values for  $S_m$  at solar elevations  $h > 42^\circ$  using mathematical extrapolation. Evidently, to correct Table 1, special measurements of daily courses of  $S_m$  for solar elevations  $h > 42^\circ$  would be necessary. Two user-friendly, simple methods for the calculation of the Atmospheric Integral Transparency Coefficient (AITC),  $p_2$ , were proposed in the end of the 1980s:

1) the method of Evnevich–Savikovskij, which may be considered consisting of two models, ES-1 and ES-2, and expressed by formulas (3) and (4) respectively, and

2) the method of Mürk–Ohvril, model MO-1, expressed by formula (6).

In order to compare the accuracy of the above-mentioned models for the calculation of AITC, a table of standard values of direct solar irradiance,  $S_m$ , proposed by Evnevich & Savikovskij (Евневич & Савиковский, 1989) was used (Table 1). According to this table, standard values of AITC,  $p_m$ , were calculated for nine different solar elevations ( $h = 10, 20, 30, ... 90^\circ$ ) and for 10 different situations of atmospheric turbidity (from very low transparency,  $p_2 = 0.410$ , to the ideal atmosphere,  $p_2 = 0.905$ ). Differences (10) between the values of  $p_m$ calculated by means of models and the standard values were considered as errors of models.

The model ES-1 demonstrated the best accuracy – the greatest error was only 2.31%. The accuracy of the model MO-1 was slightly lower, the error being 2.65%. The accuracy of model ES-2 was considerably lower – error 6.4%. So we can conclude that models ES-1 and MO-1 give almost the same accuracy, they may both be recommended for practical calculations of AITC  $p_2$ .

The advantage of the model ES-1 lies in the simplicity of formula (3), which allows easy calculation of the coefficients  $p_2$  from the results of measurements of direct solar irradiance,  $S_m$ . The advantage of the model MO-1 is that formula (6) does not contain a solar constant and is not connected with the pyrheliometric scale.

Uneven, rough angular courses of the calculated standard values of the AITC  $p_m$  (Fig. 2) prove the necessity to update the initial table of standard values of direct solar irradiance,  $S_m$  (Table 1).

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